

Silicate Atmospheres, Clouds, and Fractional Vaporization of Hot Earth-like Exoplanets

Laura Schaefer and Bruce Fegley, Jr.

Planetary Chemistry Laboratory
Department of Earth and Planetary Sciences
Washington University, St. Louis, MO

Introduction

- Exoplanets on short orbits around their stars will be hot due to both stellar and tidal heating
- Extreme heating over a planet's lifetime may lead to loss of volatiles
 - Venus is in a closer orbit than the Earth and has lost its water content due to stellar heating
 - Jupiter's moon Io is dominated by S, and may have lost lighter volatiles such as H, C, and N due to intense tidal heating
- If a hot Earth-like exoplanet loses its volatile content, a silicate atmosphere may form
- We model the formation of a silicate atmosphere of a hot Earth-like exoplanet

Super-Earth Exoplanets

- There are currently 26 known planets with $M < 20 M_{\text{Earth}}$
- Recently discovered transiting exoplanet CoRoT-7b has:
 - $M < 11 M_{\text{Earth}}$
 - $a = 0.017$ AU (orbital period of 0.854 days)
- Models (Léger et al. 2009, A&A) suggest the planet is tidally-locked
 - $T_{\text{day}} = 1800 - 2600$ K, $T_{\text{night}} \sim 150$ K
- Dayside temperature is high enough to melt and vaporize rock
 - Planet may have a partial magma ocean
- Volatiles could have been blown away from the atmosphere or condensed on the cold nightside
 - Results in a net loss of elements from the magma ocean
 - Atmosphere may be composed of rock-forming elements

Methods

- We used the MAGMA code to calculate the composition of a silicate atmosphere formed by heating of different terrestrial planet analog materials
 - calculates vaporization of systems containing Si, Mg, Fe, Ca, Al, Ti, Na, K, and O
 - As a function of temperature (1500 – 3000 K) and mass-loss by isothermal fractional vaporization
 - Results give the composition of the atmosphere and total pressure
- Fractional vaporization may simulate
 1. removal of material from dayside to nightside on CoRot-7b or
 2. Loss of material from the atmosphere
- We also calculate the composition of clouds, which may alter the atmospheric composition

5

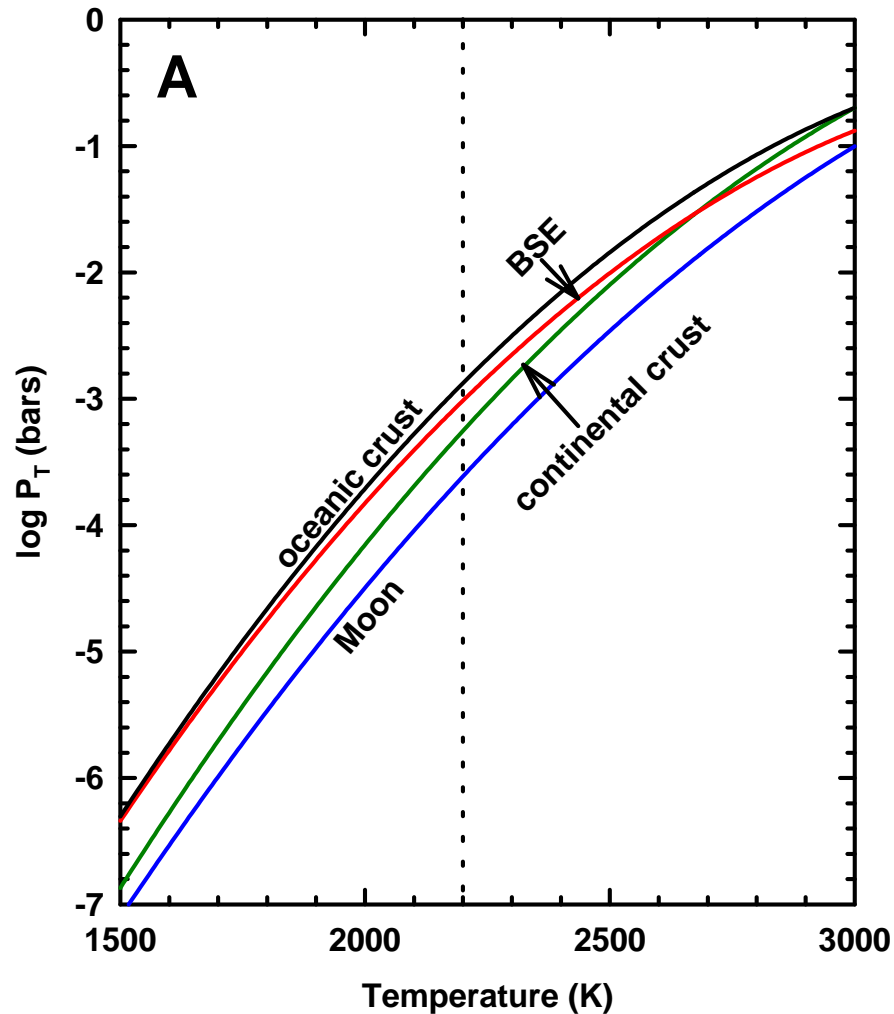
Model Planet Compositions

Oxide Wt%	Continental crust ^a	Oceanic crust ^b	BSE ^c	Moon ^d
SiO ₂	62.93	50.36	45.97	47.17
MgO	3.79	7.61	36.66	36.29
Al ₂ O ₃	15.45	15.85	4.77	3.90
TiO ₂	0.70	1.48	0.18	0.18
FeO	5.78	9.56	8.24	9.31
CaO	5.63	12.24	3.78	3.08
Na ₂ O	3.27	2.77	0.35	0.05
K ₂ O	2.45	0.13	0.04	0.005
Total	100.00	100.00	99.99	99.985

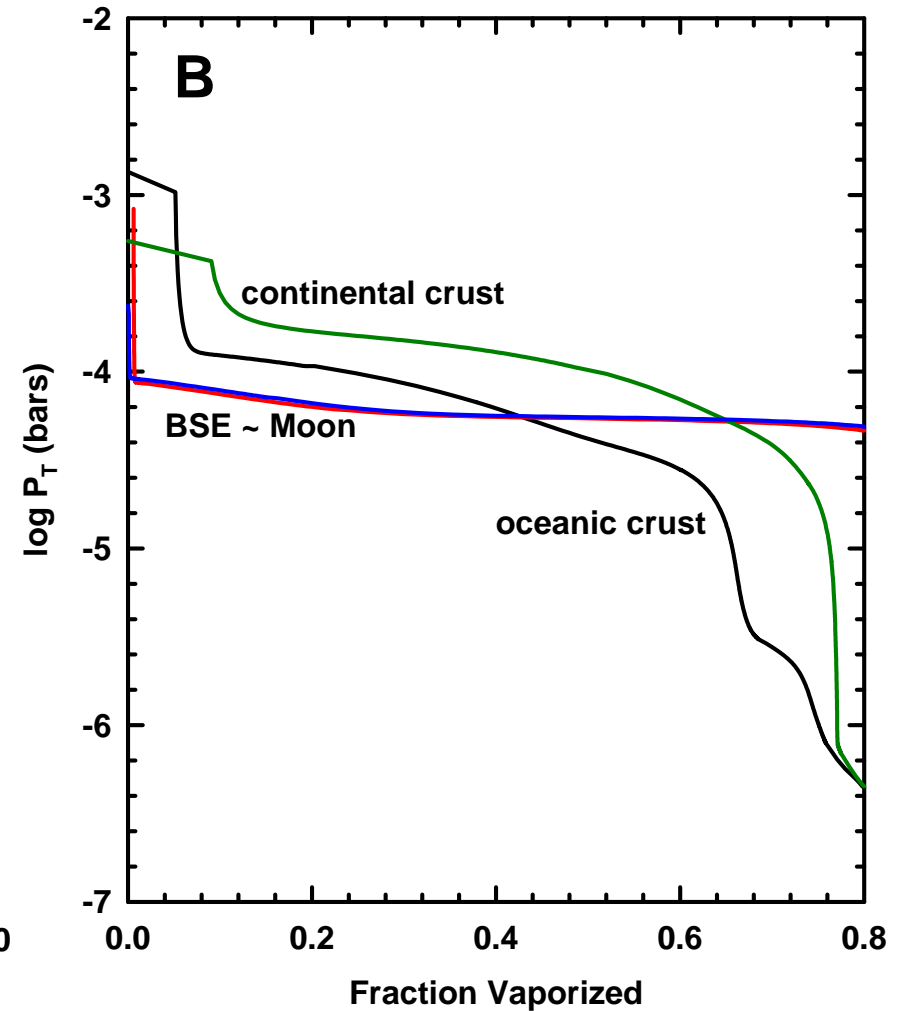
^aWedepohl (1995). ^bKlein (2005). ^cBSE=Bulk Silicate Earth. O'Neill & Palme (1998). ^dWarren (2005).

6

Total Pressure (1)



A. Initial total pressure for the model materials as a function of T (dashed line indicates $F_{\text{vap}} = 0.0$ on right graph B).



B. Total pressure for the model materials as a function of fraction vaporized at a constant temperature of 2200 K.

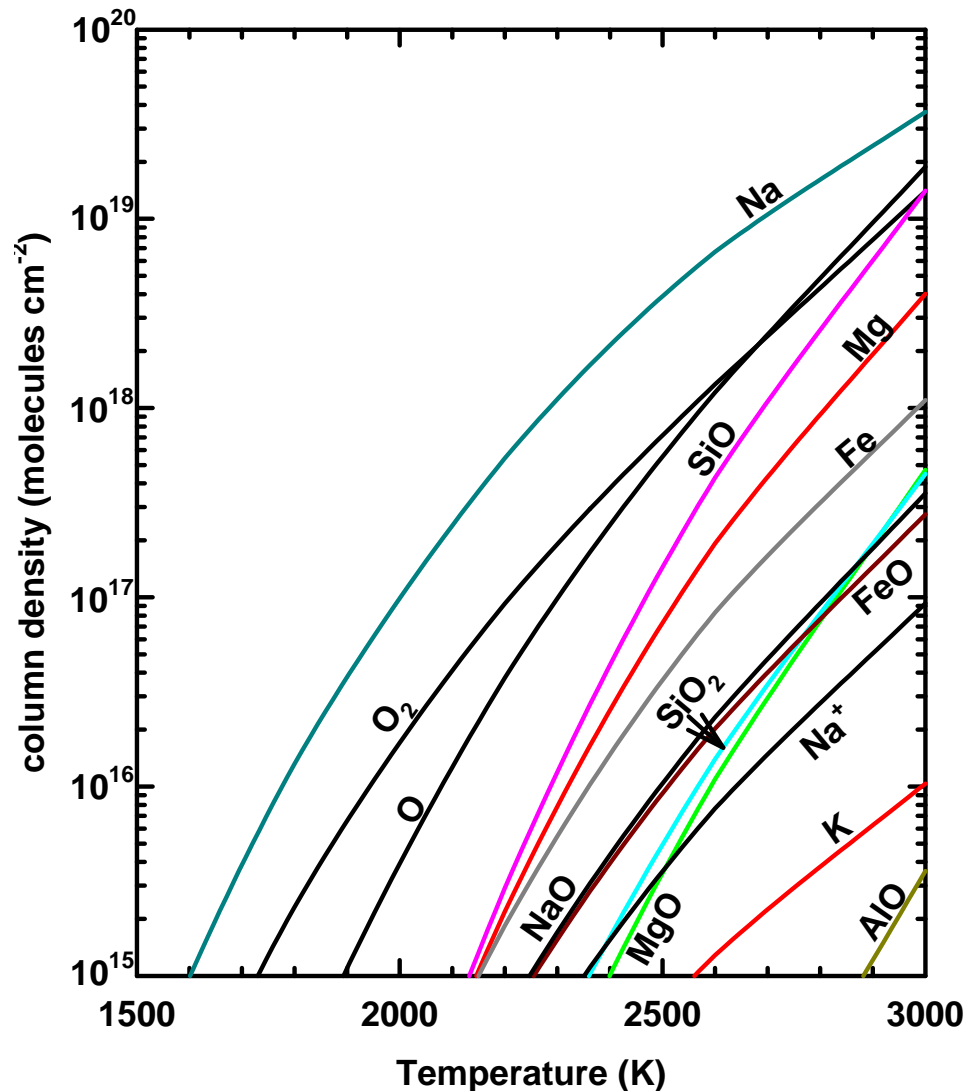
Total Pressure (2)

Material	$\log_{10} P_T = a + b/T + c/T^2$		
	<i>a</i>	<i>b</i>	<i>c</i>
Cont. crust	7.91	-29,390	10.729×10^6
Oceanic crust	6.02	-21,614	4.567×10^6
BSE	5.77	-21,524	0.494×10^6
Moon	8.17	-32,114	13.735×10^6

- Equations are for initial pressures shown in graph A on previous slide
- Initial pressures are unrelated to planet composition
- Pressure drops during fractional vaporization due to loss of material from the magma ocean

8

Temperature-dependent results

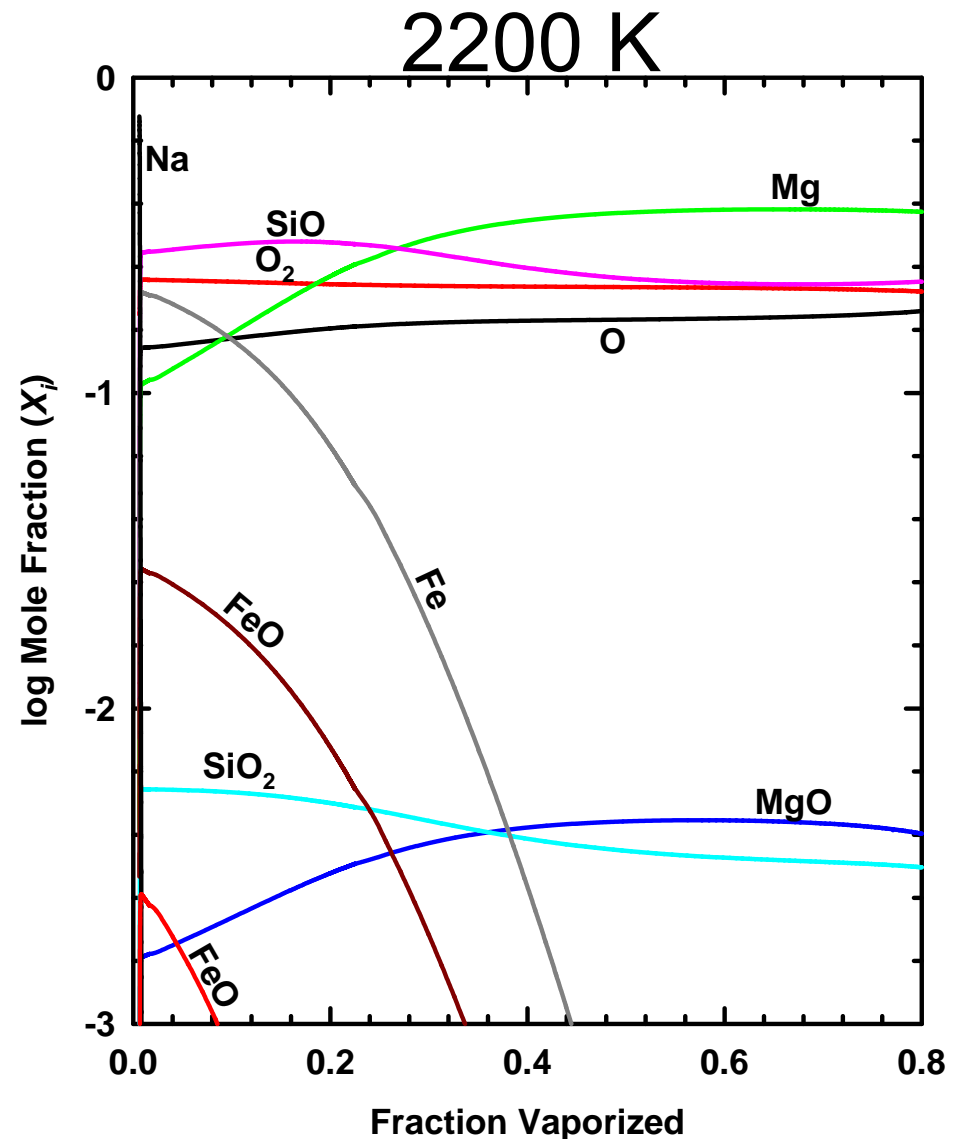


- Graph shows initial results for the BSE ($F_{\text{vap}} = 0$)
- Column density ($P_i N_A / \mu g$) is calculated for a planet the size of CoRoT-7b ($g \sim 36 \text{ m/s}^2$)
- Na is the major gas at all temperatures
 - O₂, O, and SiO are also very abundant
- Results for other model planets are similar, but:
 - K gas is much more abundant for the continental crust material
 - Fe is more abundant for the lunar material

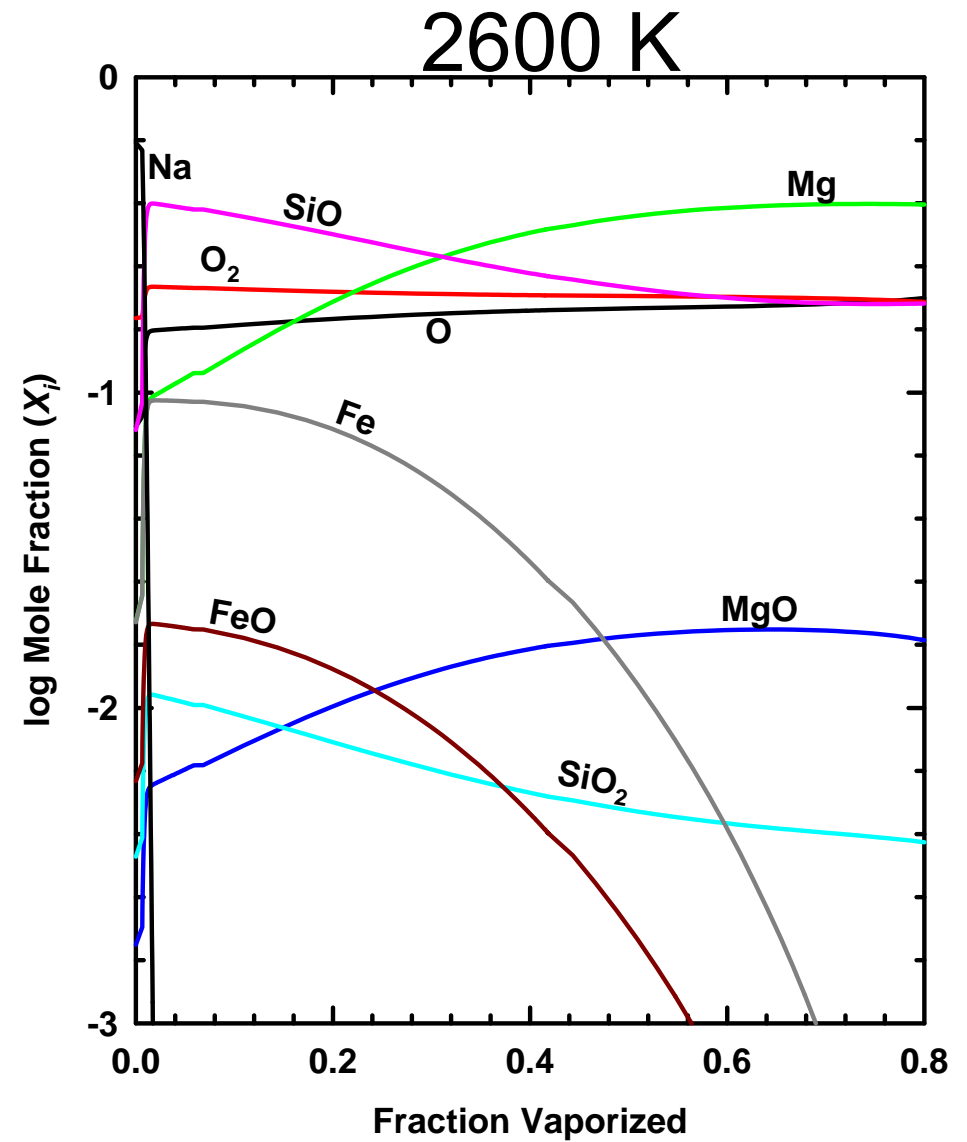
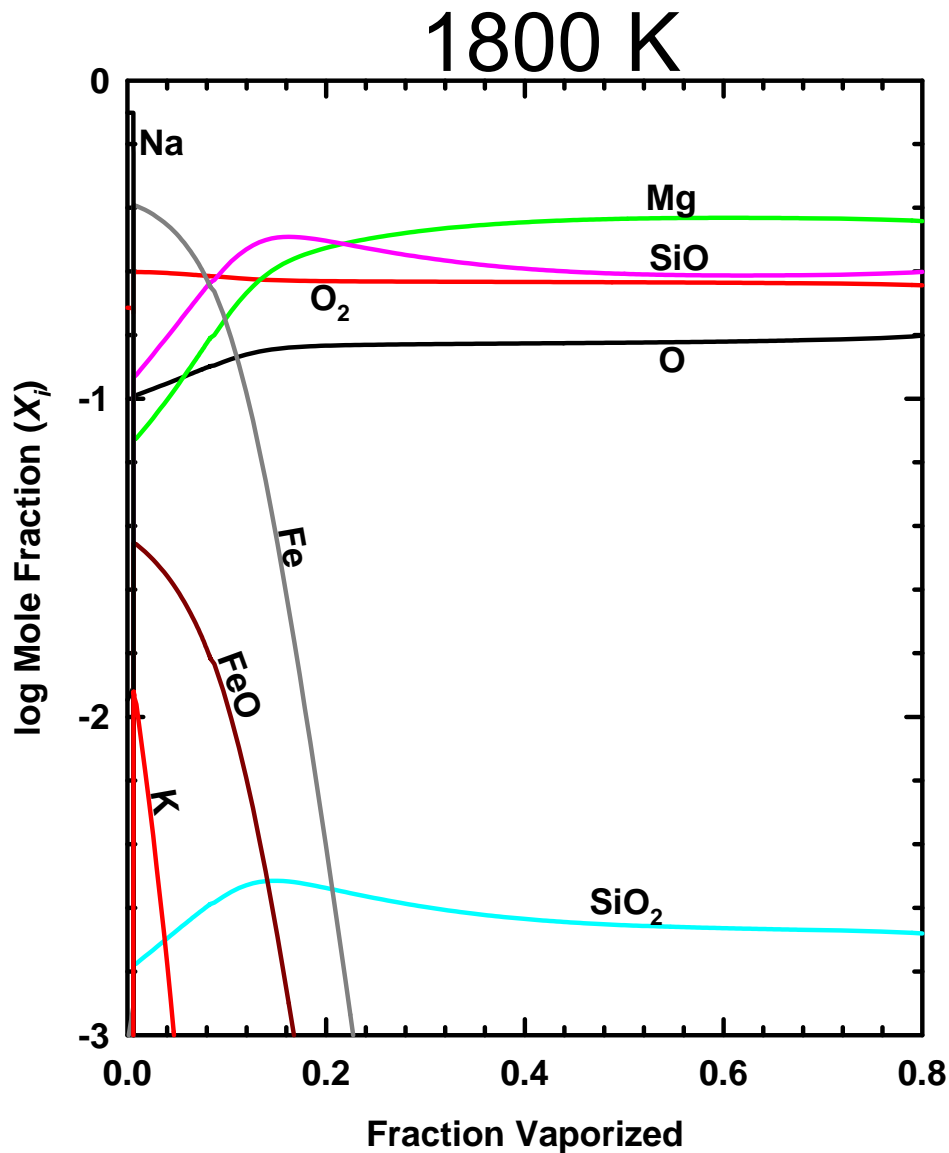
9

Fractional Vaporization of the BSE (1)

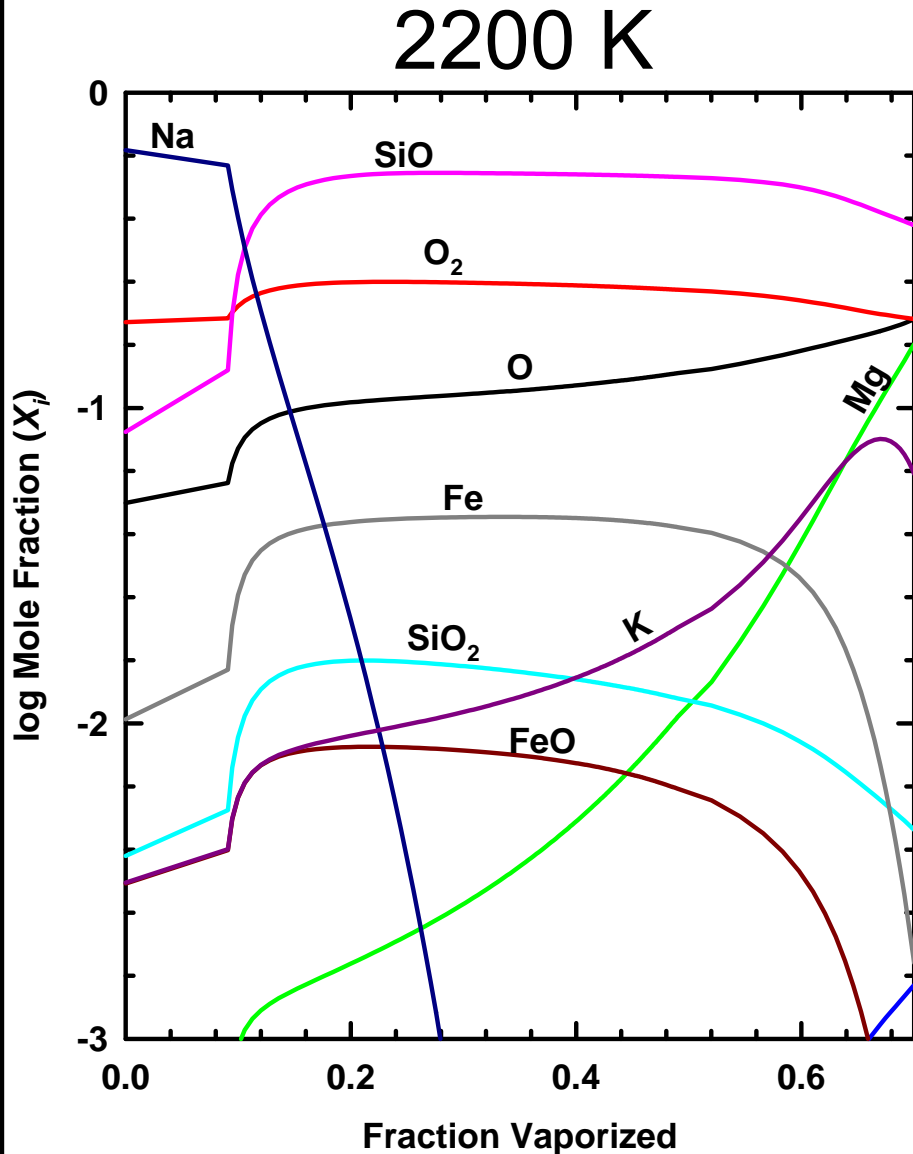
- Graphs (here and following slide) show atmospheric composition as a function of fraction vaporized at constant temperature
- Na is lost from system first, then K and Fe
- SiO becomes major gas
- Mg becomes more abundant than SiO at higher fractions vaporized
- O and O₂ maintain fairly constant abundance
- Elements are lost more quickly as temperature increases
 - e.g., see Fe(g) at 1800 K vs. 2600 K
- Results for the Moon are very similar



10 Fractional Vaporization of the BSE (2)



Continental Crust



- Graph shows atmospheric composition as a function of fraction vaporized at 2200 K
- Na is present to much higher F_{vap} than for the BSE
- SiO is major gas, and Mg is less important than for the BSE
- K is abundant at high F_{vap}
- Atmospheric composition is less dependent on F_{vap} than BSE
- Results for the oceanic crust are similar

12

Clouds in BSE atmosphere

cloud	T_{cond} (K)	Z (km)	cloud	T_{cond} (K)	Z (km)
MgSiO ₃	2125	2	“FeO”	2000	5.5
MgAl ₂ O ₄	2075	3.5	CaTiO ₃	1775	12
Al ₂ O ₃	2025	5	MgTiO ₃	1750	12.5
CaSiO ₃	2025	5	Na ₂ O	1225	27.5
SiO ₂	2025	5	K	400	50

- We calculated cloud condensation temperatures (T_{cond}) for the atmosphere generated at 2200 K for the BSE model
- Assumes dry adiabat and $g \sim 36 \text{ m/s}^2$
- Mg, Al, Si, Ca, and Fe may fall back to surface and be reincorporated in the magma ocean
- Ti, Na, and K remain in atmosphere to high enough altitudes that they may either be transported to nightside or removed from atmosphere by stellar wind

- Large clouds of Na exist around Mercury and Jupiter's moon Io
 - At Mercury, the Na cloud ($\sim 10^{11} \text{ cm}^{-2}$) extends to $\sim 23R_{\text{Mercury}}$
 - At Io, the Na cloud ($10^{10} - 10^{12} \text{ cm}^{-2}$) extends to $\sim 500R_{\text{Jupiter}}$ ($\sim 19,600R_{\text{Io}}$)
 - These clouds are very bright spectral features
- An Earth-like exoplanet with a silicate atmosphere may have an extended Na cloud
 - Na is present in the atmosphere to high altitudes
 - may interact with stellar wind
 - A large Na cloud around a transiting planet like CoRoT-7b will occult more of the stellar disk than a closely bound atmosphere
 - Increases the probability of detection for a super-Earth
- Na has already been detected in the atmospheres of several giant exoplanets (HD209458b, HD189733b)

Silicon Monoxide (SiO)

- SiO gas has strong IR bands at 4 and 9 μm and lines throughout mm wavelengths
- Possible SiO gas masers
 - SiO masers are observed in circumstellar shells, molecular clouds, star-forming regions, and supernovae
 - SiO has column densities from 10^{11} to 10^{15} cm^{-2} in these regions
 - The BSE has SiO column densities $>10^{11}$ cm^{-2} for $T > 1650$ K
- Photochemical destruction of SiO occurs by absorption of UV light ($\lambda < 300$ nm)
 - However, silicate vapor is very opaque for high temperatures and pressures, so photochemical destruction of SiO may not be very efficient
- Condensation of SiO as MgSiO_3 or SiO_2 is plausibly the most important loss process for SiO gas

Conclusions

- The major gas of silicate atmospheres formed by heating of Earth-like planets is Na(g)
 - Na stays in atmosphere to high altitudes and may interact strongly with the stellar wind
 - May form large extended clouds similar to those seen around Mercury and Io
 - Extended Na clouds would occult more of the stellar disk than a closely bound atmosphere and may be observable using current techniques
- We recommend searching for the gases Na, O₂, O, and SiO in the atmospheres of super-Earth exoplanets